

Simulating idealised satellite radiance observations

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OBJECTIVE: including simulated satellite radiance observations within a Data Assimilation scheme based on an idealised model for convection to facilitate future research

1. Introduction

An **idealised model for convection** developed at the University of Leeds and suitable for Data Assimilation research^[1,2] will be used for **satellite Data Assimilation (DA)** research. A new idealised **observation operator** has been developed in order to assimilate simulated **satellite radiance observations**. An investigation of satellite radiance DA at different scales will be performed at a later stage of the project.

2. The idealised model for convection^[1,2]

$$\begin{cases} \partial_t h + \partial_x(hu) = 0 \\ \partial_t(hu) + \partial_x(hu^2 + P) + hc_0^2 \partial_x r - \frac{1}{Ro} hv = -Q \partial_x b \\ \partial_t(hv) + \partial_x(huv) + \frac{1}{Ro} hu = 0 \\ \partial_t(hr) + \partial_x(hur) + h\tilde{\beta} \partial_x u + \alpha hr = 0 \end{cases}$$

Suitable for DA research: simpler than operational DA schemes

with:

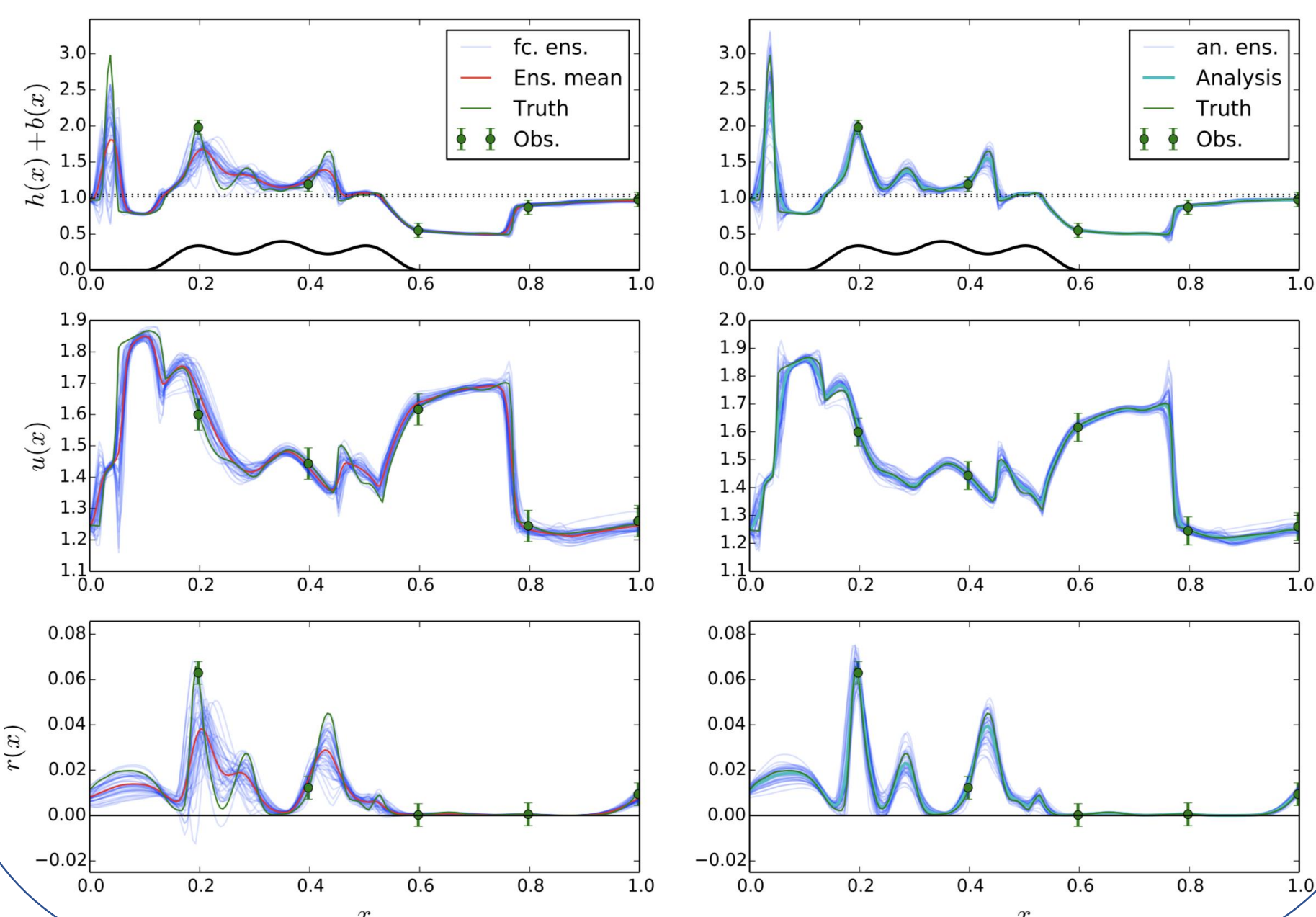
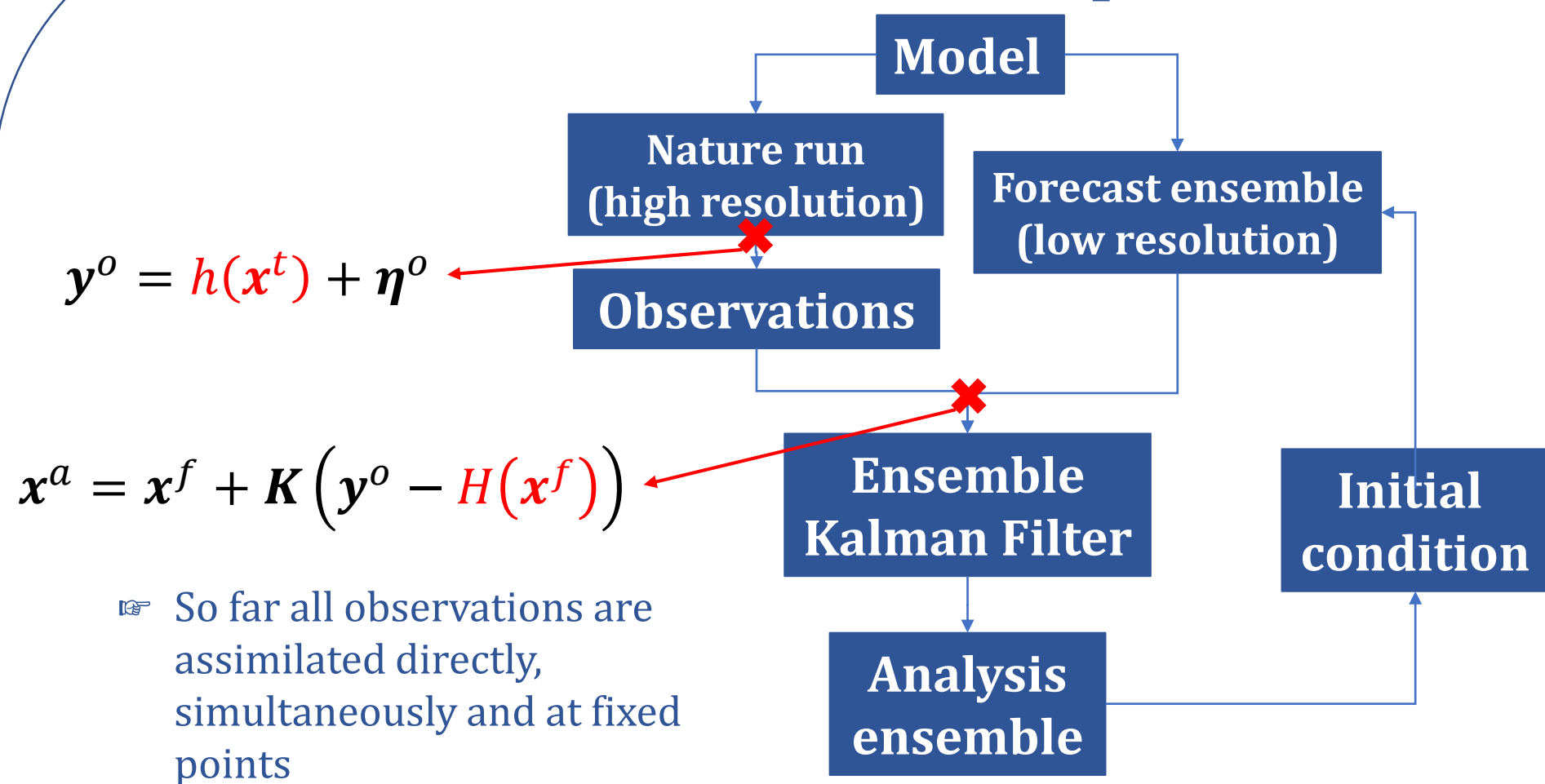
$$P = \begin{cases} \frac{1}{2}gh^2, & h + b \leq H_c \\ \frac{1}{2}g(H_c - b)^2, & h + b > H_c \end{cases}$$

$$\tilde{\beta} = \begin{cases} \beta, & h + b > H_r \text{ and } \partial_x u < 0, \\ 0, & \text{otherwise} \end{cases}$$

derivative of P

Convection and precipitation are triggered by the surface height exceeding 2 different thresholds, H_c and H_r .

3. The Data Assimilation setup^[1]



4. Simulated radiance observations

Satellite observations (emitted radiance) \longleftrightarrow ? \longleftrightarrow Model variables (surface height)

From radiance to brightness temperature at ground, Rayleigh-Jeans' law:

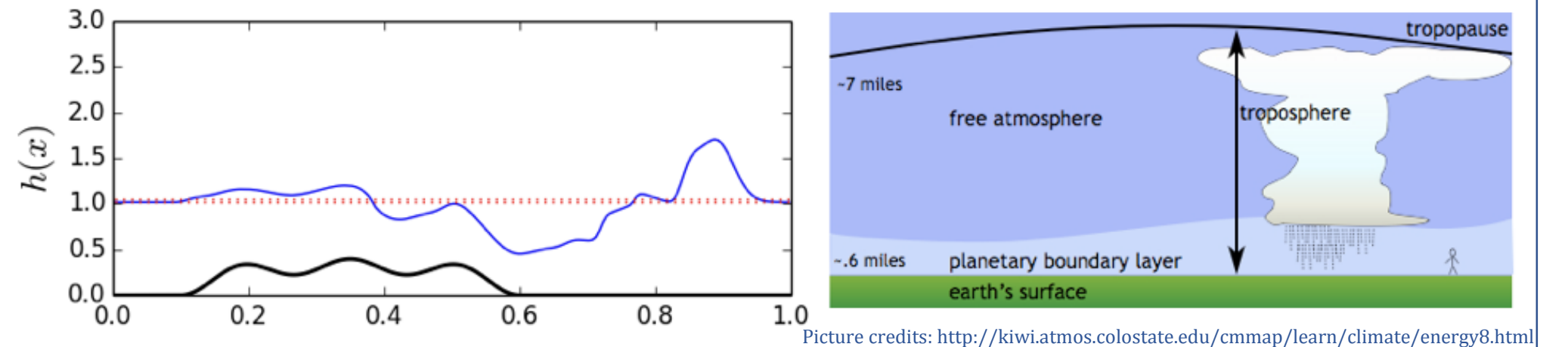
$$B(\lambda, T) = 2 \frac{k_B C}{\lambda^4} T;$$

From brightness temperature at ground to surface height, an ideal gas in hydrostatic equilibrium:

$$T = \frac{p_{top}}{\rho R} + \frac{gH}{R} h';$$

Temperature scaling T_0

Is this reasonable? Are we modelling the PBL?



$T = 284 \text{ K}$ (standard $g = 9.81 \frac{m}{s^2}$)

$T = 268 \text{ K}$ (reduced $g' = 6.3 \frac{m}{s^2}$)

Non-dimensional brightness temperature:

$$T' = \frac{T}{T_0} = \frac{p_{top}}{\rho g H} + h'$$

A less arbitrary observational variance for h' in covariance matrix $\mathbf{R} = \text{diag}(\sigma_{h'}, \sigma_{u'}, \sigma_{v'}, \sigma_{r'}) \mathbf{I}$:

$$\sigma_{h'} = \frac{1}{T_0} \sigma_T \rightarrow \text{Brightness temperature observational variance (from real instruments)}$$

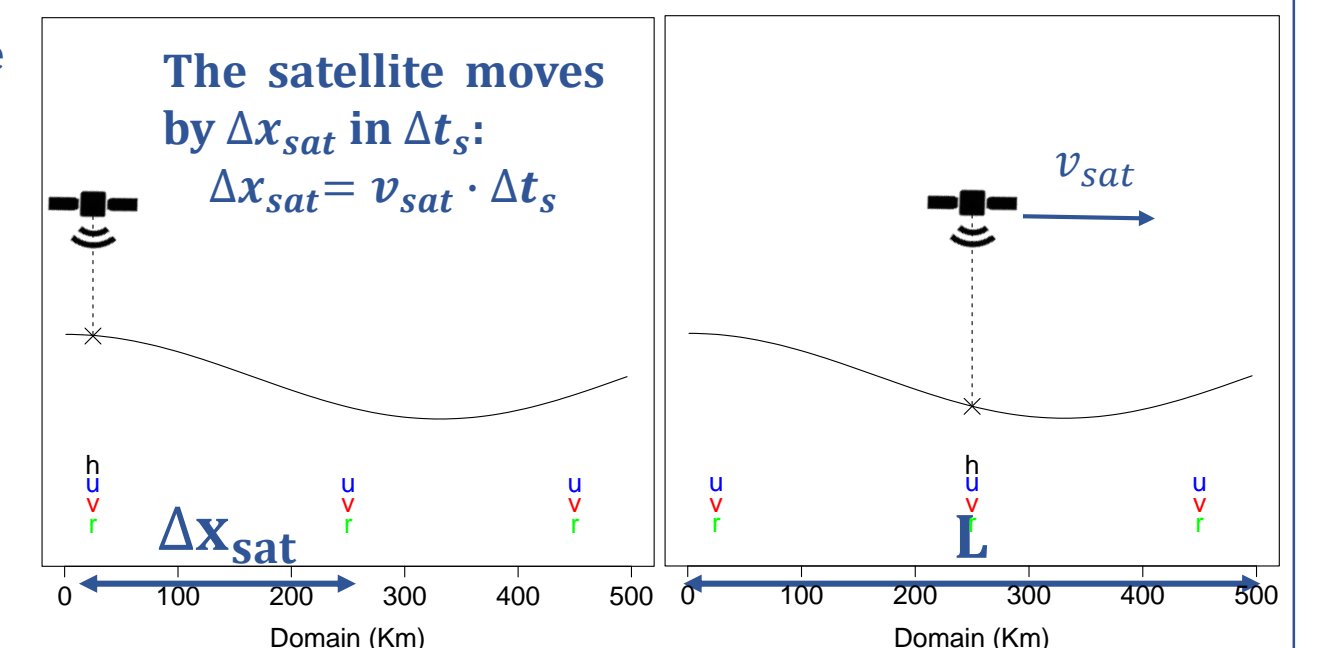
5. A spatially varying observational window

A basic configuration (periodic domain):

Δt_s time between 2 consecutive assimilation

After n assimilations:
 $t_n = \Delta t_s \cdot n$
 $x_{sat} = n \cdot \Delta x_{sat} \pmod{L}$

Observation operator:
 $H(h^f) = h(x_{sat}) + \frac{p_{top}}{\rho g H}$

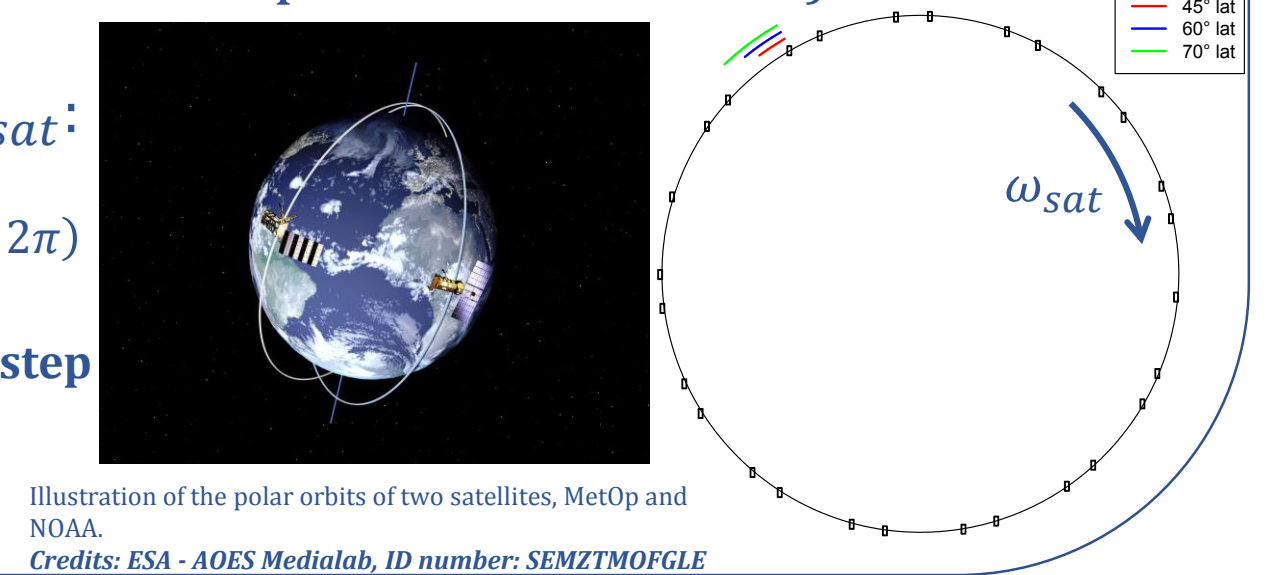


A more realistic one (based on a polar-orbit satellite):

After m orbital periods T_{sat} :

$x_{sat,1} = \omega_{sat} \cdot m \cdot T_{sat} \pmod{2\pi}$
 $x_{sat,l} = \varphi + \omega_{sat} \cdot m \cdot T_{sat} \pmod{2\pi}$

! Irregular duration of forecast step as $\Delta t_s \neq T_{sat}$



6. Future steps and questions

- Ongoing testing and tuning for the new configurations
- How can we introduce clouds?
- How can we evolve the observation operator towards something more complex and realistic?
- At a later stage: investigating the impact of satellite radiance DA at different spatial scales

References:

- [1] Kent, T. (2016): An idealised fluid model of convective-scale NWP: dynamics and data assimilation. PhD Thesis, University of Leeds.
- [2] Kent, T., Bokhove, O., Tobias, S.M. (2017): Dynamics of an idealized fluid model for investigating convective-scale data assimilation. Tellus A: Dynamic Meteorology and Oceanography, 69(1), 1369332.